

1 Reanalysis Precipitation Products over the Continental US: Are they  
2 Skillful for Century-scale Analyses?

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19 **Abstract:**

20 Within the past decade, two long-term reanalysis products (i.e. 1915-2010) have been developed  
21 to enable ‘century-scale’ analyses. These twentieth century reanalysis (20CR) products are  
22 produced by: NOAA (20CRv2), and ECWMF (ERA-20C). Using daily precipitation  
23 observations of CONUS, we study the temporal variations ranging from annual to daily scale.  
24 Our evaluation shows that 20CRv2 overestimates the precipitation magnitude, while ERA-20C  
25 shows consistent underestimation. The reconstructions of precipitation are comparable between  
26 two datasets at coarser scales (annual and monthly scales), although 20CRv2 provides better  
27 estimates at finer temporal scales (weekly and daily scales). Both 20CRs correctly capture the  
28 precipitation variations in the west coast of US, and these variations are well captured at daily  
29 scale in 20CRv2 data. These results suggest the potential of using 20CR products to reconstruct  
30 storms events in the west coast area during 1915-1948 when there are currently no other  
31 reanalysis products available.

32 **Keywords:** reanalysis, precipitation, climate, numerical modeling.

## 33 1. Introduction

34 Twentieth century reanalysis (20CR) product is the model reconstruction of climate in the  
35 past century. They are built for climatological study, and they have demonstrated skill in  
36 capturing key climate variations such as North Atlantic Oscillation and El Nino [*Giese et al.*,  
37 2010; *Compo et al.*, 2011]. They have also been used to evaluate other numerical simulations  
38 such as GCMs [*Sheffield et al.*, 2013]. Previous studies have shown their ability to capture the  
39 long-term statistics and variations in precipitation at annual to monthly scales in different regions  
40 [*Ferguson and Villarini*, 2012; *Misra et al.*, 2013; *Zhang et al.*, 2013; *Kong and Bi*, 2015].  
41 Efforts have been made to evaluate the precipitation products in 20CR, but all of them have  
42 restricted the analysis to monthly scale due to the limitation of observed data. For example,  
43 *Ferguson and Villarini* (2012) evaluated the temperature and precipitation simulation in the  
44 central US using Climate Research Unit (CRU) monthly observation and concluded that the  
45 20CR produced by NOAA overestimates the precipitation amount when compared with CRU  
46 observation. However, the quality of 20CR in event-scale simulations and over longer time  
47 period (such as a century) has not yet been investigated.

48 Modern atmospheric numerical models require initial and boundary conditions as model  
49 input, which are often obtained from atmospheric reanalysis datasets. Several reanalysis datasets  
50 have been developed for the post-1948 atmospheric conditions [*Kalnay et al.*, 1996; *Kanamitsu*  
51 *et al.*, 2002; *Mesinger et al.*, 2006]. However, 20CR remains the only source of boundary  
52 conditions for numerical modeling of events before 1948. Recent studies suggest that although in  
53 general 20CR is not a good option for precipitation event reconstruction, it may be usable in  
54 certain climate regions such as the west coast of US [*Chen and Hossain*, 2016]. From a modeling  
55 perspective, historical event reconstruction can also be viewed as dynamic downscaling of the

56 reanalysis data used for boundary conditions. The difficulty of precipitation simulation has been  
57 recognized as a result of highly variable rain rates and large spatial variability  
58 [*Intergovernmental Panel on Climate Change et al.*, 2013]. Thus, it is possible that in  
59 areas/durations where reanalysis precipitation shows a good match with observation at coarser  
60 resolution, numerical model downscaling may produce skillful results. Therefore, it is important  
61 and useful to identify the areas/periods where the 20CR can serve for such event-scale  
62 reconstructions. Using quality-controlled ground-based precipitation observations as reference,  
63 this study aims to examine the spatial-temporal precipitation reconstruction quality over the  
64 continental US (CONUS) in the currently available 20CRs. The goal is to understand how  
65 skillful the currently available reanalysis products are for analyses of precipitation events over  
66 the last century (1915-2010).

## 67 **2. Data and Analysis**

68         Currently there are two 20CRs available: one is the Twentieth Century Reanalysis  
69 Product (20CRv2) produced through a collaboration between NOAA and University of Colorado  
70 Cooperative Institute for Research and Environmental Sciences; the other (ERA-20C) is  
71 produced by European Centre for Medium-Range Weather Forecasting (ECMWF), and it is an  
72 outcome of the ERA-CLIM project. Table 1 shows the details of these two 20CRs. The 20CRv2  
73 is an ensemble product (56 ensemble members), and in this study only the mean values of  
74 ensembles are used.

75         The evaluation of the products used the Livneh daily CONUS near-surface gridded  
76 meteorological dataset [*Livneh et al.*, 2013] as the reference observation. This precipitation  
77 dataset was generated from rain gauge records since 1915. Given its long-term temporal  
78 coverage, it has been used in other model evaluation studies [*Sheffield et al.*, 2013]. In this study,

79 the data over CONUS (see Figure 1) was used, and all the data were first conservatively  
80 regrided to the global Gaussian T62 grids, the coarsest resolution among these three datasets,  
81 before the analysis was conducted. Variations in the reconstructed precipitation were evaluated  
82 using the spatial and temporal correlation between 20CR precipitation and Livneh precipitation  
83 datasets.

### 84 **3. Results and Discussion**

85 Figure 1 shows the climatology of precipitation during 1915-2010. The pattern of higher  
86 annual precipitation in both west and east coast area in the reference data is correctly captured in  
87 both reconstructions. Regarding the multiyear average annual precipitation amount, 20CRv2  
88 shows a better match in the southeast area, though both failed to give out the observed large  
89 amount of precipitation in the Pacific Northwest area. Figure 1 also suggests that 20CRv2 tends  
90 to overestimate the precipitation in the center and the east US. ERA-20C provides a better  
91 matched spatial pattern, but it slightly underestimates the precipitation in the southeast US.

92 To understand the reconstructed spatial distribution of annual precipitation, annual spatial  
93 correlation coefficients were calculated. Figure 2 shows these correlations along with the factors  
94 that would influence these correlations. It is clear that the correlation coefficients (thus  
95 reconstruction of annual precipitation maps) are superior in skill in recent years, as one would  
96 expect. The two products not only agree with each other but also yield high correlation  
97 coefficients for the recent period of 1980-2010. As we go back in the past, such as pre-1950, the  
98 20CRv2 outperforms ERA-20C. Also, for the wet years (larger circles in Figure 2), 20CRs tend  
99 to produce more accurate annual precipitation maps.

100 Figure 3 shows the temporal correlation of seasonal and monthly precipitation series. The  
101 large patterns are similar at seasonal and monthly scales, with the highest skill observed in the  
102 west coast, followed by the mid-west. For these two regions, two 20CRs show similar  
103 performance, with ERA-20C slightly better in central north. In the southeast area (i.e. Florida),  
104 however, 20CRv2 clearly outperforms the ERA-20C. The climate of the west coast can partly  
105 explain the high correlation in this region. Previous studies have pointed out that in this region  
106 (especially California), over 60% of annual precipitation is contributed by less than 20 events  
107 [Sun *et al.*, 2006], which are often the atmospheric rivers events [Hagos *et al.*, 2016].  
108 Atmospheric river events are one of the large-scale meteorological events that numerical models  
109 can capture better than the more localized convective storms.

110 It is also important to note that the reanalysis products are capable of describing the  
111 monthly cycle of precipitations. Here we focus on the average monthly cycle, and all the  
112 monthly data were taken to calculate multiyear averaged monthly precipitation and standard  
113 deviation. The results are shown in Figure 4, and it is clear that for all 12 months, 20CRv2  
114 usually overestimate, while ERA-20C tends to underestimate the precipitation amount. In the  
115 months except June-September, the precipitation in ERA-20C matches closer with the  
116 observation. In summer time, however, 20CRv2 gives out better estimates. This confirms the  
117 superiority of 20CR in capturing heavy rainfall events. Previous studies have suggested that  
118 taking 20CR and other reanalyses as “ensemble” could improve the reconstruction quality, such  
119 as the surface mass balance of Greenland ice sheet [Hanna *et al.*, 2011]. The comparison in  
120 Figure 4 also suggests that combining the information from 20CRv2 and ERA-20C may produce  
121 a better estimate on precipitation patterns.

122 To investigate the skill of the 20CRs to produce correct spatial-temporal structures in  
123 precipitation, the monthly precipitation data are decomposed using empirical orthogonal  
124 functions (EOF). The analysis was done on the detrended monthly data, in which the linear  
125 trends at each grid were taken out. The results of EOF analysis are shown in Figure 5. Panels (a-c)  
126 are the first mode in monthly precipitation, panel (d) shows the contribution of each mode in the  
127 monthly variation. Both 20CRv2 and ERA-20C produce similar first modes as seen in the  
128 observation: the transition of opposite phases from west to east coast. However, the first mode in  
129 20CRv2 is closer to the observation, where the phase in the southeast inland area is less  
130 significant. ERA-20C overestimates the area of negative phase in the west US. Regarding the  
131 magnitude of variation, panel (d) shows that the contribution of the first mode in 20CRv2 is  
132 almost the same as that from observation, but ERA-20C gives out better estimates for the next a  
133 few modes (mode 2 to 5).

134 Figure 6 shows the spatial distribution of correlations on weekly and daily precipitation  
135 series. At weekly scale, 20CRv2 and ERA-20C show similar performance, with 20CRv2 being  
136 slightly more skillful in the eastern US. At the daily scale, however, 20CRv2 shows much better  
137 performance. In both datasets, the daily variations in the west coast are significantly better  
138 constructed, with 20CRv2 showing better overall correlation coefficients. The modeling efforts  
139 on several extreme rainstorms before 1948 also suggests that 20CRv2 can be used to reconstruct  
140 the west coast storms reasonably well, but the reconstruction at other areas are heavily biased  
141 [*Chen and Hossain, 2016*]. It is also important to note that the correlation of daily precipitation  
142 has different trends in two 20CRs, as shown in Figure 7. In 20CRv2, the correlation remains  
143 between 0.55-0.60 for the whole duration of 1915-2010, and the correlation is slightly higher in  
144 recent years. ERA-20C data shows a peak of correlation around 1960-1975, and the

145 reconstruction in this period is comparable to 20CRv2. However, the correlation after 1975 has a  
146 systematic decrease, indicating three distinct epochs regarding the skill of daily precipitation: 1)  
147 1915-1960, when the reconstruction shows the lowest quality; 2) 1960-1975, when the  
148 reconstruction is best; 3) 1975-2010, when the reconstruction quality is of medium quality.

149         Although 20CRs were originally built for climate reconstruction, our analysis suggests  
150 that it may also be suitable for event-scale precipitation simulations over the West Coast of US.  
151 This finding can potentially encourage science/engineering communities to extend analyses with  
152 confidence further back than the 1950s to re-evaluate storms of the early 20<sup>th</sup> century. For  
153 example, the Hydrometeorological Reports (HMR) published by NOAA [*Schreiner and Riedel,*  
154 1978] has outlined a collection of extreme historical rainstorms for engineering safety design  
155 purposes. Previous studies have shown the capability of reanalysis products in rebuilding those  
156 rainstorms after 1948 using various numerical models [*Tan, 2010; Woldemichael et al., 2014;*  
157 *Chen and Hossain, 2016*]. However, a significant portion of this HMR collection occurred before  
158 1948, for which direct meteorological observations other than precipitation are limited.  
159 Therefore, numerical modeling is the only viable approach to obtain a more complete physical  
160 picture about these events. The good quality in 20CRs at least over the west coast suggests that  
161 we can now apply numerical modeling for extreme precipitation events that took place in the  
162 first half of the 20<sup>th</sup> century.

#### 163 **4. Conclusions**

164         In this study we employed a gauge-based precipitation dataset to evaluate the  
165 precipitation simulated by two 20CR during 1915-2010 at various temporal scales. The major  
166 findings are:



- 167 (1) Both 20CRs show strengths in building the precipitation climatology, at scales from  
168 yearly to monthly. However, there is significant decrease in the reconstruction quality  
169 of weekly/daily precipitation in the whole CONUS;
- 170 (2) The skill of precipitation reconstruction varies in different parts of CONUS, with the  
171 west coast indicating good quality even in daily scale. Thus it is possible to use 20CR  
172 for the heavy storms simulation in the west coast area;
- 173 (3) 20CRv2 is better at capturing the statistics in heavy rainfall events/periods, while  
174 ERA-20C is better at describing light rainy periods;
- 175 (4) 20CRv2 tends to overestimate the precipitation climatology, while ERA-20C shows  
176 consistent underestimations. Using both 20CRs as “ensemble” for quantitative  
177 precipitation reconstruction is worthwhile.

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179 **Data Policy:** All data are available from public sources. Specifically, 20CRv2 data was obtained  
180 from UCAR (University Corporation for Atmospheric Research) Research Data Archive at  
181 <http://rda.ucar.edu> as ds131.2 dataset; ERA-20C data was obtained from the same website as  
182 ds626.0 dataset. Livneh precipitation database was obtained from NOAA's Earth System  
183 Research Laboratory at <http://www.esrl.noaa.gov/psd/data/gridded/data.livneh.html>. The data  
184 that have been processed for the analysis are available upon request to the first author.

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## References

187 Chen, X., and F. Hossain (2016), Revisiting Extreme Storms of the Past 100 Years for Future  
188 Safety of Large Water Management Infrastructures, *Earth's Future*,

189       doi:10.1002/2016EF000368.

190   Compo, G. P. et al. (2011), The Twentieth Century Reanalysis Project, *Q. J. R. Meteorol. Soc.*,  
191       *137*(654), 1–28, doi:10.1002/qj.776.

192   European Centre for Medium-Range Weather Forecasts (2014), ERA-20C Project (ECMWF  
193       Atmospheric Reanalysis of the 20th Century), <http://dx.doi.org/10.5065/D6VQ30QG>,  
194       Research Data Archive at the National Center for Atmospheric Research, Computational  
195       and Information Systems Laboratory, Boulder, Colo. (Updated daily.) Accessed 16 Mar  
196       2016.

197   Ferguson, C. R., and G. Villarini (2012), Detecting inhomogeneities in the Twentieth Century  
198       Reanalysis over the central United States, *J. Geophys. Res. Atmos.*, *117*(5), 1–11,  
199       doi:10.1029/2011JD016988.

200   Giese, B. S., G. P. Compo, N. C. Slowey, P. D. Sardeshmukh, J. A. Carton, S. Ray, and J. S.  
201       Whitaker (2010), The 1918/19 El Niño, *Bull. Am. Meteorol. Soc.*, *91*(2), 177–183,  
202       doi:10.1175/2009BAMS2903.1.

203   Gilbert P. Compo et al. (2015), NOAA/CIRES Twentieth Century Global Reanalysis Version 2c,  
204       <http://dx.doi.org/10.5065/D6N877TW>, Research Data Archive at the National Center for  
205       Atmospheric Research, Computational and Information Systems Laboratory, Boulder, Colo.  
206       (Updated yearly.) Accessed 16 Mar 2016.

207   Hagos, S. M., L. R. Leung, J. Yoon, J. Lu, and Y. Gao (2016), A projection of changes in  
208       landfalling atmospheric river frequency and extreme precipitation over western North  
209       America from the Large Ensemble CESM simulations, *Geophys. Res. Lett.*, *43*, 1–7,  
210       doi:10.1002/2015GL067392.

211 Hanna, E. et al. (2011), Greenland Ice Sheet surface mass balance 1870 to 2010 based on  
212 Twentieth Century Reanalysis, and links with global climate forcing, *J. Geophys. Res.*  
213 *Atmos.*, 116(24), 1–20, doi:10.1029/2011JD016387.

214 Intergovernmental Panel on Climate Change et al. (2013), *Climate Change 2013 - The Physical*  
215 *Science Basis*.

216 Kalnay, E. et al. (1996), The NCEP/NCAR 40-year reanalysis project, *Bull. Am. Meteorol. Soc.*,  
217 77(3), 437–471, doi:10.1175/1520-0477(1996)077<0437:TNYRP>2.0.CO;2.

218 Kanamitsu, M., W. Ebisuzaki, J. Woollen, S. K. Yang, J. J. Hnilo, M. Fiorino, and G. L. Potter  
219 (2002), NCEP-DOE AMIP-II reanalysis (R-2), *Bull. Am. Meteorol. Soc.*, 83(11), 1631–  
220 1643+1559, doi:10.1175/BAMS-83-11-1631.

221 Kong, X., and X. Bi (2015), Dynamical Downscaling of the Twentieth Century Reanalysis for  
222 China: Climatic Means during 1981–2010, *Atmos. Ocean. Sci. Lett.*, 8(3), 166–173,  
223 doi:10.3878/AOSL20140099.

224 Livneh, B., E. A. Rosenberg, C. Lin, B. Nijssen, V. Mishra, K. M. Andreadis, E. P. Maurer, and  
225 D. P. Lettenmaier (2013), A long-term hydrologically based dataset of land surface fluxes  
226 and states for the conterminous United States: Update and extensions, *J. Clim.*, 26(23),  
227 9384–9392, doi:10.1175/JCLI-D-12-00508.1.

228 Mesinger, F. et al. (2006), North American regional reanalysis, *Bull. Am. Meteorol. Soc.*, 87(3),  
229 343–360, doi:10.1175/BAMS-87-3-343.

230 Misra, V., S. M. DiNapoli, and S. Bastola (2013), Dynamic downscaling of the twentieth-century  
231 reanalysis over the southeastern United States, *Reg. Environ. Chang.*, 13(1), 15–23,

232 doi:10.1007/s10113-012-0372-8.

233 Schreiner, L. C., and J. T. Riedel (1978), *Probable maximum precipitation estimates, United*  
234 *States east of the 105th meridian*, Department of Commerce, National Oceanic and  
235 Atmospheric Administration.

236 Sheffield, J. et al. (2013), North American Climate in CMIP5 experiments. Part I: Evaluation of  
237 historical simulations of continental and regional climatology, *J. Clim.*, 26(23), 9209–9245,  
238 doi:10.1175/JCLI-D-12-00592.1.

239 Sun, Y., S. Solomon, A. Dai, and R. W. Portmann (2006), How often does it rain?, *J. Clim.*,  
240 19(6), 916–934, doi:10.1175/JCLI3672.1.

241 Tan, E. (2010), Development of a Methodology for Probable Maximum Precipitation Estimation  
242 over the American River Watershed Using the WRF Model, University of California, Davis.

243 Woldemichael, A. T., F. Hossain, and R. Pielke (2014), Impacts of Postdam Land Use/Land  
244 Cover Changes on Modification of Extreme Precipitation in Contrasting Hydroclimate and  
245 Terrain Features, *J. Hydrometeorol.*, 15(2), 777–800, doi:10.1175/jhm-d-13-085.1.

246 Zhang, Q., H. Körnich, and K. Holmgren (2013), How well do reanalyses represent the southern  
247 African precipitation?, *Clim. Dyn.*, 40(3-4), 951–962, doi:10.1007/s00382-012-1423-z.

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## Tables

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Table 1. Details of two 20CRs

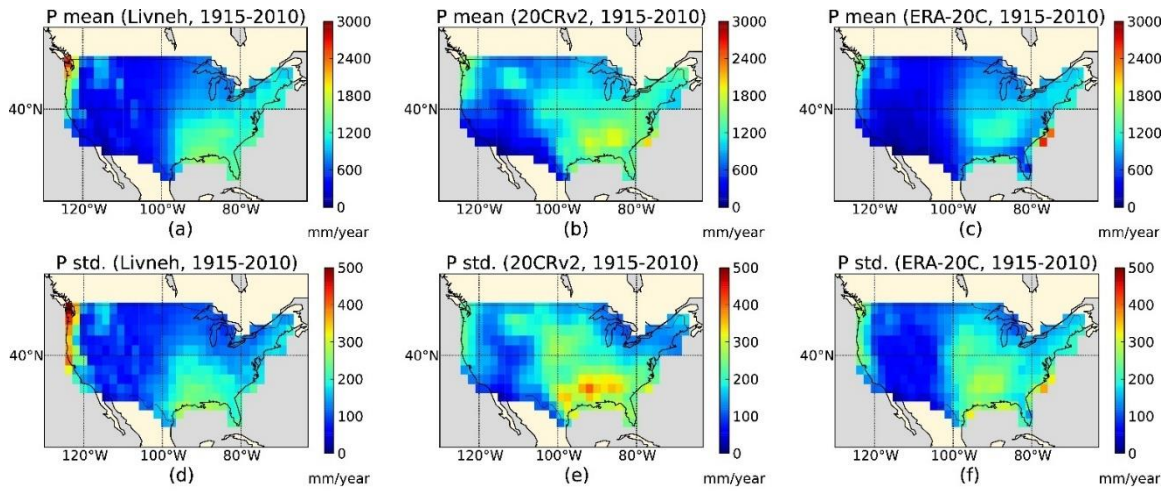
Item	20CRv2	ERA-20C
Produced by	NOAA-CIRES	ECMWF
Temporal coverage	1850/12/31-2014/12/31	1900/1/1-2010/1/1
Temporal resolution	6-hourly	3-hourly
Spatial coverage	Global	Global (89.142S – 89.142N)
Spatial resolution	T62 grid (~210km)	T159 grid (~125km)
Vertical layers	28	91
Data assimilated	Surface observations of synoptic pressure	Surface and mean sea level pressure; Surface marine winds
Data assimilation method	Ensemble Kalman filter	4D-VAR

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## Figures

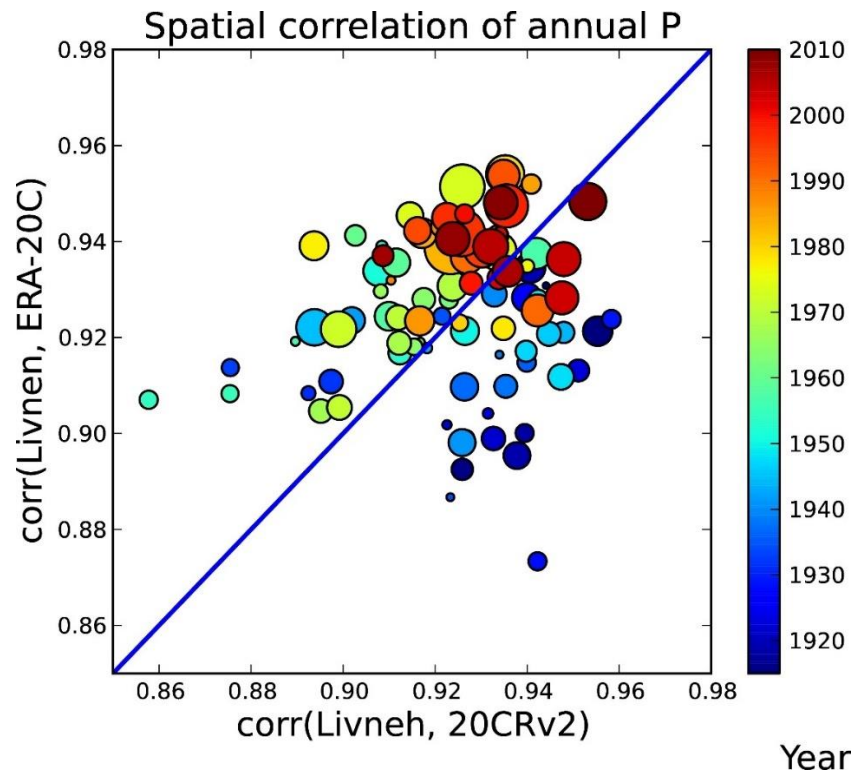


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256 **Figure 1. Precipitation climatology (1915-2010) from observation and model reconstructions. (a)**  
257 **multiyear average annual precipitation in the reference dataset; (b) and (c) multiyear average**  
258 **annual precipitation in 20CRv2 and ERA-20C. Panels (d-f) are the standard deviation of annual**  
259 **precipitation.**

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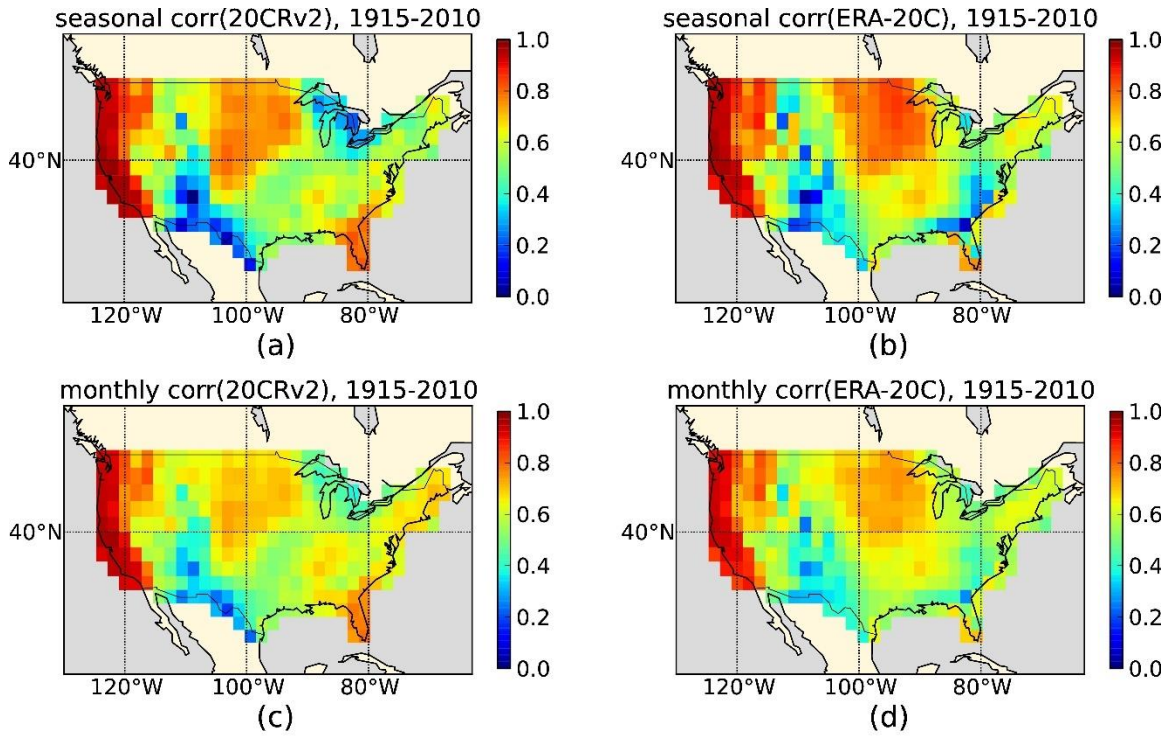
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263 **Figure 2. Spatial correlation between annual precipitation maps. Colors of points indicate different**  
264 **years, and sizes of points are the total rain from the given year (from Livneh dataset).**

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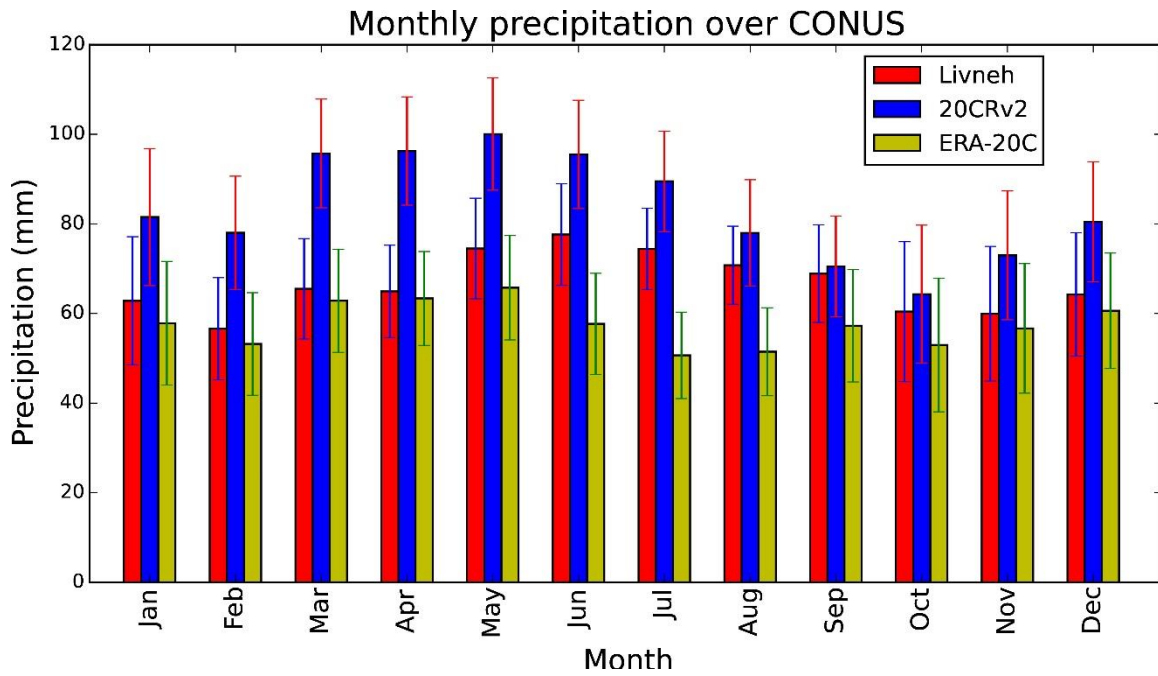
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**Figure 3. Spatial distribution of correlation of seasonal and monthly precipitation.**

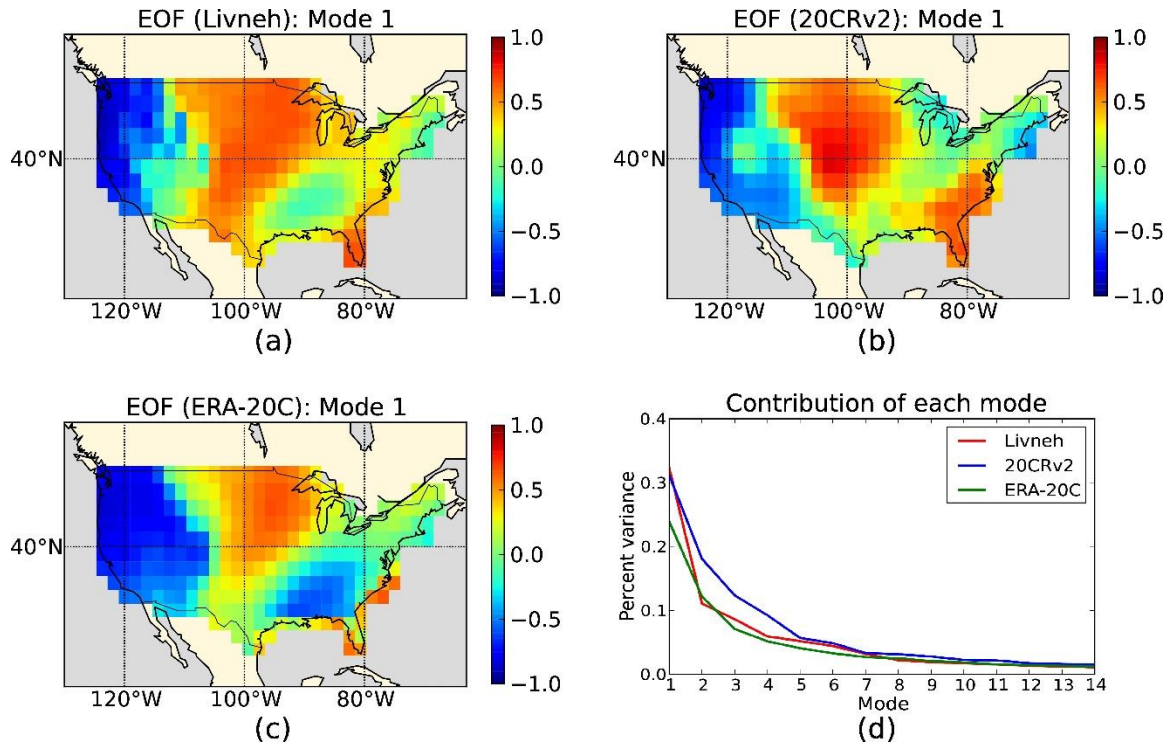




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270 **Figure 4. Statistics of monthly precipitation over CONUS. Error bars show the standard deviation.**

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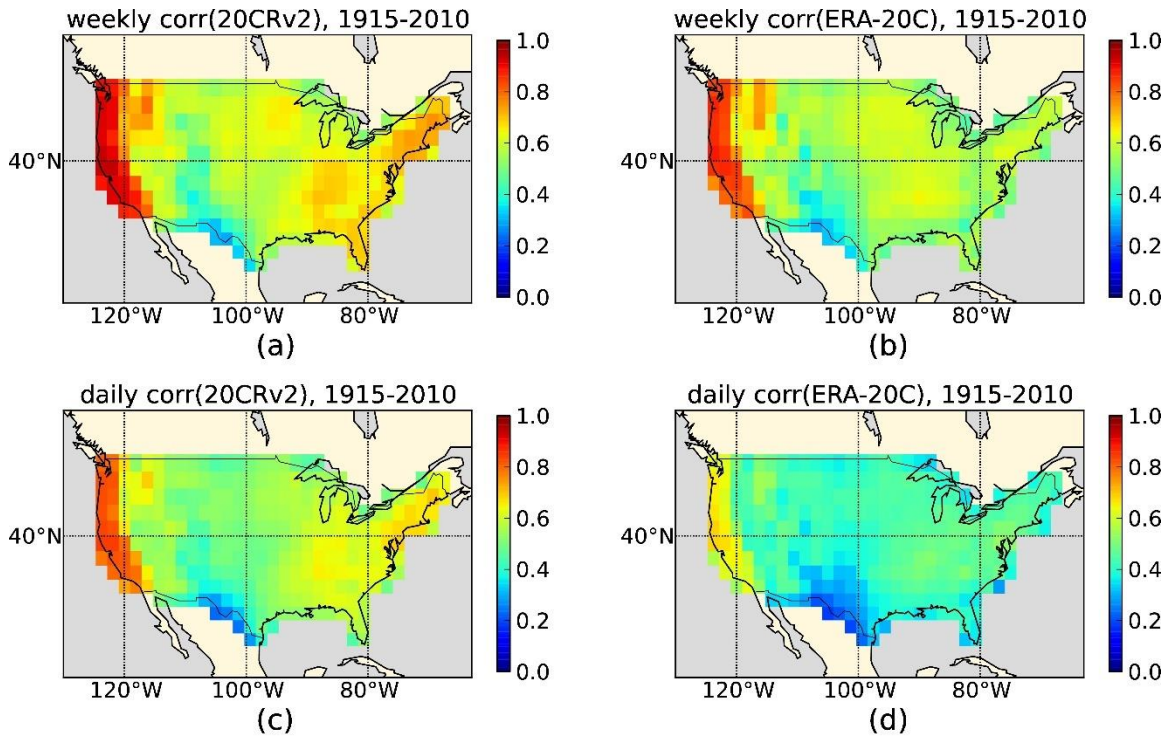


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**Figure 5. EOF analysis of monthly precipitation data.**

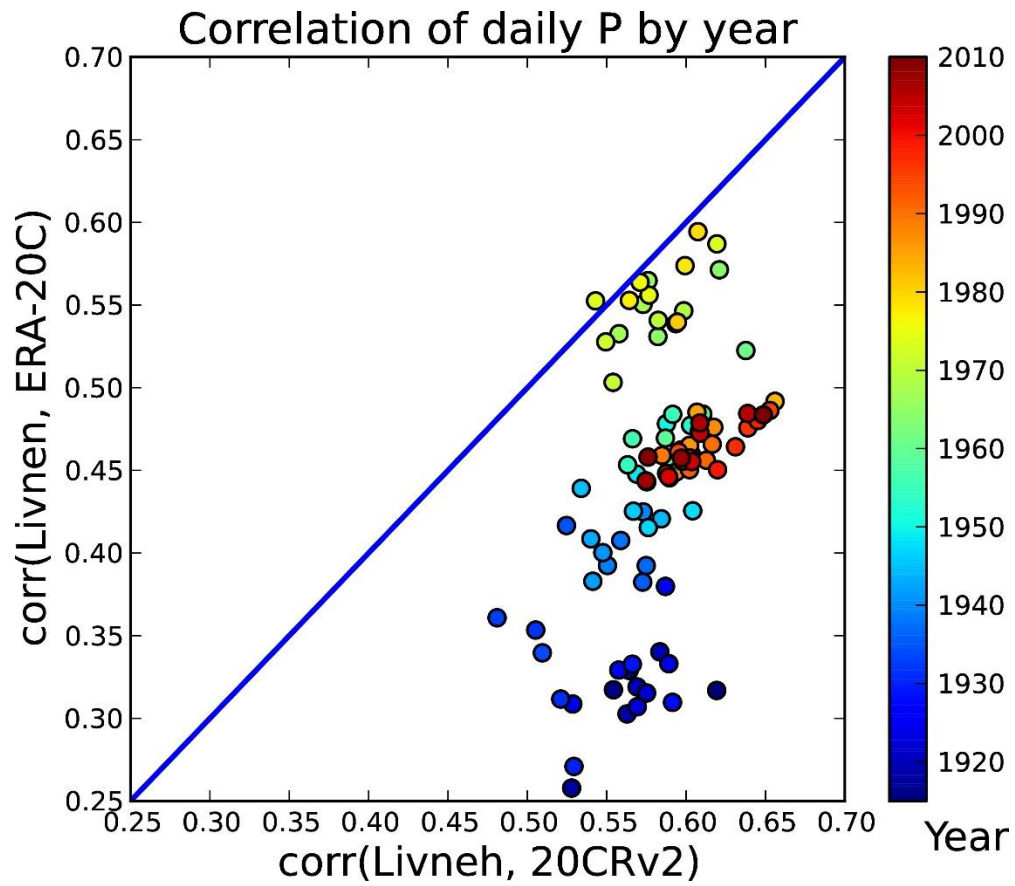


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**Figure 6. Spatial distribution of correlation of weekly and daily precipitation.**



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**Figure 7. Correlation of daily precipitation as a function of year.**